Accepted Manuscript

Effects of moisture content on wind erosion thresholds of biochar

F.C. Silva, C. Borrego, J.J. Keizer, J.H. Amorim, F.G.A. Verheijen

PII: S1352-2310(15)30487-8

DOI: 10.1016/j.atmosenv.2015.10.070

Reference: AEA 14228

To appear in: Atmospheric Environment

Received Date: 30 April 2015

Revised Date: 23 October 2015

Accepted Date: 26 October 2015

Please cite this article as: Silva, F.C., Borrego, C., Keizer, J.J., Amorim, J.H., Verheijen, F.G.A., Effects of moisture content on wind erosion thresholds of biochar, *Atmospheric Environment* (2015), doi: 10.1016/j.atmosenv.2015.10.070.

This is a PDF file of an unedited manuscript that has been accepted for publication. As a service to our customers we are providing this early version of the manuscript. The manuscript will undergo copyediting, typesetting, and review of the resulting proof before it is published in its final form. Please note that during the production process errors may be discovered which could affect the content, and all legal disclaimers that apply to the journal pertain.





Adhesion forces between small particles increase resistance to wind erosion (>15% moisture content)

1 Effects of moisture content on wind erosion thresholds of biochar

- 2 F.C. Silva¹, C. Borrego¹, J.J. Keizer¹, J.H. Amorim^{1,2}, F.G.A. Verheijen^{1,*}
- 3
- 4 ¹University of Aveiro, Centre for Environmental and Marine Studies (CESAM), Department of Environment
- 5 and Planning, Campus Santiago, 3810-193 Aveiro, Portugal

6 ²Swedish Meteorological and Hydrological Institute (SMHI), Air quality research unit, SE-60176 Norrköping,

7 Sweden

- 8 *Corresponding author: <u>frankverheijen@gmail.com</u>, Tel.: +351 234 370 200
- 9
- 10

11 Highlights

- Wind erosion of biochar was assessed through wind tunnel simulations
- 13 Moisture content lower than 10 % does not prevent erosion of fine particles
- Higher moisture increases adhesion of fine particles and weight of large particles
- 15 Minimum 15 % moisture is recommended for reducing wind erosion of biochar
- 16

17 Abstract

Biochar, i.e. pyrolysed biomass, as a soil conditioner is gaining increasing attention in research and industry, 18 19 with guidelines and certifications being developed for biochar production, storage and handling, as well as for 20 application to soils. Adding water to biochar aims to reduce its susceptibility to become air-borne during and 21 after the application to soils, thereby preventing, amongst others, human health issues from inhalation. The 22 Bagnold model has previously been modified to explain the threshold friction velocity of coal particles at 23 different moisture contents, by adding an adhesive effect. However, it is unknown if this model also works for 24 biochar particles. We measured the threshold friction velocities of a range of biochar particles (woody 25 feedstock) under a range of moisture contents by using a wind tunnel, and tested the performance of the 26 modified Bagnold model. Results showed that the threshold friction velocity can be significantly increased by 27 keeping the gravimetric moisture content at or above 15 % or greater to promote adhesive effects between the 28 small particles. For the specific biochar of this study, the modified Bagnold model accurately estimated 29 threshold friction velocities of biochar particles up to moisture contents of 10 %.

30

31 Keywords

32 biochar; threshold friction velocity; wind tunnel; particle size; Bagnold model

33 Abbreviations

| ρ | Density |
|---------|---|
| β | Adhesive effect parameter |
| Α | Aerodynamic constant |
| a | Conceptual contribution of gravity to threshold friction velocity |
| В | Regression coefficients of the response surface methodology |
| b | Conceptual contribution of adhesive forces to threshold friction velocity |
| d_m | Mean particle size diameter |
| dP | Pressure differential |
| k | von Kármán constant |
| RSM | Response surface methodology |
| U | Free stream velocity |
| U^* | Threshold friction (or shear) velocity |
| W | Gravimetric water content |
| Ζ | Height |
| | |

34

49

35 1. Introduction

Airborne dust particles have raised concerns for both environmental (Choobari et al., 2014; IPCC, 2013) and
human health reasons (De Capitani et al., 2007; Hashizume et al., 2010; Karanasiou et al., 2012; Salvi and
Barnes, 2009). Known sources of dust aerosols, including black carbon, range from biomass and fossil fuel
burning (Andreae and Merlet, 2001; Ito and Penner, 2005; Jacobson, 2001), wind erosion of aggregate storage
piles of coal and charcoal (Toraño et al., 2007), to rail and road transport, including through re-suspension
(Buchsbaum, 2007; Ferreira et al., 2003; Harrison et al., 2012)

42 The concept of biochar, where biomass is pyrolised to create a char that can improve soil functioning

43 (Lehmann, 2007), has received increasing scientific attention in the last years (Verheijen et al., 2014). The

44 main research focus has been on soil carbon sequestration (Lehmann et al., 2006; Nguyen and Lehmann,

45 2009) and crop yields (Jeffery et al., 2011), but biochar has also been observed to change many other soil

46 processes and functions (Lehmann and Joseph, 2015). The pyrolysis process can cause carbonaceous dust

47 emissions, although pyrolysis generally produces substantially less black carbon aerosols than biomass

48 burning (Whitman et al., 2011). Nonetheless, biochar particles can also become air-borne after the pyrolysis

process, namely: i) during post-processing, packing, storage and transport; ii) during application in the field;

50 and iii) through soil erosion by wind during the lifetime of biochar in soil. Considering the post-production

51 process, the European Biochar Certificate recommends that biochar should be kept "sufficiently moist to

52 prevent dust generation or dust explosions", but does not quantify the relevant moisture contents (Schmidt et

al., 2012). At the start of writing this paper, a search of the SCOPUS database for "biochar and wind" only

54 retrieved one relevant article, i.e. a paper about applying biochar below the soil surface to avoid the risk of

erosion by wind (Blackwell et al., 2010). Verheijen et al. (2010) also stated that there is a paucity of data on

56 interactions between biochar and wind to inform sustainable biochar development.

Relationships between moisture content and threshold friction velocities have been observed for coal dust 57 58 (Duo-min and Shu-tang, 1991) and have also been modelled for a range of coal particle sizes by means of the modified Bagnold model (Zhang et al., 2012; 2013). However, biochar has different physical properties than 59 60 coal and it is unknown if the model developed for coal particles is also valid for biochar particles. The 61 objectives of this study were, therefore: i) to determine the relationships between moisture content and 62 threshold friction velocity for the full range of particles ($<50 \mu m$ to $>6,000 \mu m$) of a common, woody 63 feedstock biochar; ii) to test if the contribution of adhesive and gravity forces can be predicted by the 64 modified Bagnold model; and iii) to discuss the implications of our findings for biochar production and application. To this end, we conducted wind erosion experiments in the wind tunnel of the Department of 65 Environment and Planning (University of Aveiro) to determine threshold velocities for biochar particles of 6 66 67 size classes and at 6 moisture contents.

68

69 **2. Materials and methods**

70 2.1. Biochar source and preparation

Biochar was purchased from Swiss-Biochar GmbH, where it was produced from a mixed wood sievings feedstock in a Pyreg[®] 500 III pyrolysis unit, 620 °C maximum temperature, 20 min duration, 80 % C content and H/C ratio 0.18. The main physico-chemical characteristics are presented in Table 1. Six particle size classes were obtained by mechanical sieving at 5000, 3150, 2000, 200 and 50 µm mesh widths. The resulting particle size classes are identified according to their mean particle size as shown in Table 2, where the largest particle size class was assumed to have a mean of 6,000 µm. Subsequently, the biochar fractions were ovendried at 75 °C for 48 h.

- 78 The different biochar particle size classes were combined with six contents of gravimetric moisture to achieve
- a full-factorial experiment of two factors and six levels per factor, with 3 to 5 replicates for each of the 36
- combinations. The gravimetric water contents (1, 3, 6, 10, 15 and 20 %) were selected to cover a similar range
 as used in other studies on the influence of moisture content on particle transport by wind (Chen et al., 1996;
- 82 Zhang et al., 2012).
- The moisture contents were attained by calculating the required mass of water, assuming that the density of
 distilled water is approximately 1 kg L⁻¹, and verified by measuring weight loss following heating at 105 °C
 for 24 h. Specific volumes of biochar were put into plastic bags and then wetted with a liquid dispenser
- 86 (atomizer) producing a fine spray (Cornelis and Gabriels, 2003; Han et al., 2009). The plastic bags were then
- 87 placed in a refrigerator at 4° C for a minimum period of 24 h to achieve homogeneous wetting of the biochar.
- 00
- 88
- 89

90 Table 1. Bulk biochar physico-chemical characteristics.

| Characteristic | Value (average ± standard deviation) |
|--|--------------------------------------|
| Electrical conductivity of the leachate (μ S cm ⁻¹) | $1,496 \pm 43$ |
| pH | 8.13 ± 0.04 |
| pH of the leachate | 9.96 ± 0.07 |
| Ash content (%) | 10.36 ± 0.97 |
| Density (kg m ⁻³) | 184 ± 4 |

91

92 **2.2.** Wind tunnel experiments

93 2.2.1. Experimental setup

94 The experimental work was performed in the wind tunnel laboratory facilities of the Department of

95 Environment and Planning, at the University of Aveiro, Portugal. These facilities consist of an open-circuit,

96 suction type wind tunnel, with a test section of 7×1.5×1 m (length × width × height), as described in (Borrego
97 et al., 2007).

Biochar samples were placed over a white surface (for maximizing colour contrast) that was centred on the
 test section floor of the wind tunnel. A square of 36 cm² was filled with sample material up to a height of

approximately 0.5 cm (Figure 1). During the actual wind tunnel experiments, three moments were defined for

101 visual observation and recording of the pressure differential (dP): (i) moment 1, when sample particles started

to vibrate, corresponding to the threshold shear stress; (ii) moment 2, when particles from the sample edges

103 started to be transported, representing the incipient motion; and (iii) moment 3, when particles from the

sample centre started to be transported, representing the major contribution of the entrainment (Dey, 1999).

105 More specifically, moments 2 and 3 were defined as when five particles had been eroded. As in several other

106 wind tunnel soil erosion experiments (e.g., Alfaro et al., 1997; Dong et al., 2003; He et al., 2008; Genis et al.,

107 2013), a Pitot tube was used for the measurement of wind velocity, mainly due to its robustness when working

108 with particle-fed flows. Free stream velocity ranged from 0.63 to 10.28 m s^{-1} , with a rotor frequency-step

109 between consecutive measurements of 2.5 Hz (corresponding to an average velocity-step of 0.54 m s⁻¹). The

110 vertical wind profile was determined by measuring dP at 12 heights (Z) from the wind tunnel floor, ranging

from 0.5 to 50.0 cm (the latter corresponding to half the height of the wind tunnel). Free stream wind velocity

112 ranged from 0.0 to 10.3 m s⁻¹.

113

102



114

- 115 Figure 1. Biochar sample on wind tunnel test section floor.
- 116

117 2.2.2. Experimental determination of threshold friction velocity

- 118 The vertical wind velocity profile in the atmospheric boundary layer under neutral stability conditions can be
- 119 described by the logarithmic relation:

$$U(Z) = \left(\frac{U^*}{k}\right) ln\left(\frac{Z}{Z_0}\right) \tag{1}$$

120

121 , where U is the free stream velocity, U^* is the threshold friction velocity, k is the von Kármán constant (0.4), 122 and Z_0 is the aerodynamic roughness height.

123 Eq. (1) may be approximated by a least-squares curve fitting method (Dong et al., 2003), which, then, enables 124 the determination of U^* , if the variation of velocity with height is known. This method was applied, using the 125 values for the velocity determined by the Pitot tube and the respective height, and is represented by:

$$U(Z) = M \cdot ln(Z) + N \tag{2}$$

126

- 127 , where *M* and *N* are regression constants.
- 128 The threshold friction velocity is obtained by:

$$U^* = k \times M \tag{3}$$

129

130 **2.2.3. Statistical analysis**

The analysis of variance (two-way ANOVA with interaction model) was applied to the U* values of moment 3 (particles detaching from the centre of the sample, see Section 2.2.1), as these are the most representative of field conditions. This analysis was used to test significance of particle size and moisture content, as well as their interaction, on threshold friction velocities, and it was performed in the software package IBM SPSSTM version 20.

Furthermore, the response surface methodology (RSM) was also applied in order to better illustrate the
combined influence of particle size and moisture content. This methodology was only applied to the U* values
of moment 3, as these are the most representative of field conditions. The experimental results were modelled
according to Eq. (4):

$$U *_{mod} = B_0 + B_1 x_1 + B_2 x_2 + B_{1,2} x_1 x_2 + B_{1,1} x_1^2 + B_{2,2} x_2^2$$
(4)

140

141 , where U_{mod}^* is the response variable, x_1 represents the gravimetric water content (*W*, in %), x_2 represents the 142 mean size of each class of biochar particles (d_m , in µm), B_0 is the model constant, B_1 and B_2 are linear 143 coefficients (main effects), $B_{1,2}$ is a cross-product coefficient (interaction) and $B_{1,1}$ and $B_{2,2}$ are quadratic 144 coefficients (Myers et al., 2009). Fitting was done with the software package MatlabTM version R2011b, using 145 the least squares method. To this end, the ranges of both factors (*W* and d_m) were codified (x_1 and x_2) between 146 -1 and 1, as shown in Table 2 The goodness-of-fit was assessed by R² as well as the significance of the 147 regression model (*F*-test).

148

Table 2. Codification of the gravimetric water content (x_1) and the mean size of each class of biochar particles

150 (x_2) , for application of the response surface methodology.

| W (%) | 1 | 3 | 6 | 10 | 15 | 20 |
|------------------|-------|--------|--------|--------|--------|----|
| x_1 (codified) | -1 | -0.789 | -0.473 | -0.053 | 0.474 | 1 |
| d_m (µm) | 6,000 | 4,075 | 2,575 | 1,100 | 125 | 40 |
| x_2 (codified) | 1 | 0.354 | -0.149 | -0.644 | -0.971 | -1 |
| | | | | | | |

151

152 2.3. Modelling of the threshold friction velocity

Modelling of U^* , like their statistical analysis, was done just for the experimental data of moment 3. For each particle size class, U^* was modelled as a linear function of W, following Eq.(5):

$$U^* = bW + a \tag{5}$$

155 , where b and a represent the contributions of the adhesive forces between particles and of gravity,

- respectively. This linear relationship allows assessing which of these two forces are dominant in the threshold
- 157 friction velocity for a given particle size class (Zhang et al., 2012).
- 158 Aeolian transport of particles can furthermore be modelled as a function of the aerodynamic drag force that
- 159 lifts the particles, taking into account particle size and density as well as air density. This is usually done
- assuming that the aerodynamic related constant (*A*) has a value around 0.10 (Bagnold, 1941). The original
- 161 Bagnold model, however, did not explicitly consider the effect of the particles' water content. Therefore,
- 162 Zhang et al. (2012) recently proposed to introduce the effect of the water content, in particular for modelling

the Aeolian transport of coal particles transported by wind. This modified Bagnold's model is defined as Eq.(6):

$$U^* = A_{\sqrt{\frac{1}{1-W}}} \sqrt{\frac{\rho_{biochar}}{\rho_{air}}} g. d_m \tag{6}$$

165 , where $\rho_{biochar}$ and ρ_{air} are the densities of biochar and air, respectively, and g is the acceleration of gravity. 166 While the A in the modified Bagnold model needs to be determined experimentally, the model can be further

167 extended to determine the contribution of the adhesive forces by adding a term (β) that is defined as a power

168 law function of both gravimetric water content and mean particle diameter:

$$\beta = \frac{W^c}{d_m^d}$$

169 , where c and d are non-negative parameters.

170 The extended model then becomes Eq. (8):

$$U^* = A \sqrt{\frac{1}{1 - W}} \sqrt{\frac{W^c}{d_m^d}} \sqrt{\frac{\rho_{biochar}}{\rho_{air}}} g.d_m \tag{8}$$

171

172 **3. Results**

173 Although the data set is slightly non-linearly distributed (p-value = 0.050 for Kolmogorov-Smirnov test), 174 error's variances are equal (p-value = 0.049 for Levene test). Therefore, two-way analysis of variance was 175 performed, and showed that each of the studied factors (d_m and W), as well as their interaction on U^* , were all 176 significant at a significance level of 0.05 (Table S1 provided as supplementary material). These results 177 anticipated that the effect of particle size on threshold friction velocity is more pronounced than moisture or 178 even than the interaction between both factors. Subsequent results are, therefore, presented by focusing on the 179 effect of each separated factor on threshold friction velocity of biochar particles.

180

181 **3.1. Effect of particle size**

- 182 The observed U^* values for moments 1 and 2 (when the particles started to vibrate and to be eroded from the 183 sample edges, respectively) are given as supplementary material, in Figures S1 and S2 respectively.
- 184 The U^* values for moment 1 could not be measured for the two smaller particle size classes (d_m of 40 and 125
- 185 μ m), as indicated by the zero values in Figure S1. The observed U* values for moment 1 ranged from 0.17 m
- 186 $s^{-1} (d_m = 1,100 \ \mu m \text{ with } W = 10\%)$ to 0.33 m $s^{-1} (d_m = 4,075 \ \mu m \text{ with } W = 20\%)$. This wide range can be
- 187 attributed primarily to the heterogeneity in the size of the biochar particles. These differences in particle size

- 188 may have been associated to differences in other relevant factors such as the particles' aerodynamic 189 behaviour, shape sphericity, tortuosity, ratio between length and thickness (platyness sensu (Boton et al., 190 2013)), pore size distribution, mechanical strength, orientation, and density (including packing fraction). The 191 U^* values for moment 1 revealed a clear tendency to increase with increasing particle size (at least up to 4,075 192 μ m), regardless of the particles' moisture content. This increase in U* values was most pronounced between 193 particle sizes 2,575 to 4,075 μ m, especially for the higher moisture contents (6 – 20 %). In contrast, U* values 194 hardly changed or even decreased between particle sizes 4,075 µm to 6,000 µm, possibly because the smaller 195 particles were more platy than the larger ones, as observed by (Terzaghi, 1996), and, thus, experienced a 196 stronger shear stress.
- 197 The U^* values for moment 2, (Figure S2), showed a similar behaviour as moment 1: generally, U^* increased
- with increasing particle size, up to a d_m of 4,075 µm. Moment 2 U^* values ranged from 0.15 to 0.49 m s⁻¹, for $d_m = 125 \mu m$ at W = 6 % and $d_m = 4,075 \mu m$ at W = 15 %, respectively. Edge erosion may not represent fieldscale applications of biochar in soil. All the centre particle erosion events (moment 3) presented higher U^* values. In this study, centre particles eroding are seen as a more representative assessment of the field-scale reality and are analysed in more detail below.
- 203 The U* values for the moment 3, (Figure 2) ranged from 0.25 m s⁻¹ ($d_m = 125 \ \mu m$ at $W = 1 \ \%$) to 0.61 m s⁻¹
- 204 $(d_m = 6,000 \ \mu\text{m} \text{ at } W = 20 \ \%)$. U* values increased with increasing particle size class (except for $d_m = 6,000$ 205 $\mu\text{m} \text{ at } W = 3 \ \%$ and $W = 10 \ \%$). The most obvious pattern was that U* tended to be higher for smaller particle 206 size classes (d_m of 40 and 125 μm) when moisture contents were greater than 1 %. For d_m of 40 and 125 μm
- 207 particle size classes, 3 % and 10 % moisture contents resulted in U^* values between 0.27 and 0.30 m s⁻¹, very
- similar to those obtained for larger particles with $d_m = 1,100 \,\mu\text{m}$ (U* between 0.29 and 0.30 m s⁻¹). Further
- increases in moisture content (15 % and 20 %) resulted in significant (*p*-value < 0.001) U* increases for the
- small particle size classes ($0.38 0.39 \text{ m s}^{-1}$ and $0.41 0.43 \text{ m s}^{-1}$ for d_m of 40 and 125 μ m, respectively). This
- 211 observation is in accordance with (Zhang et al., 2012), who found that the moisture content had considerably
- greater impacts on U^* of smaller than larger coal particles. These authors reported that below $d_m = 250 \,\mu\text{m}$,
- and above W = 6 %, U^* was not particle size dependent but rather dependent on moisture content, which they
- attributed to the weight-moisture gain and inter-particle aggregation.
- 215



Figure 2. Threshold friction velocities (U*) regarding particle detachment from the sample centre (moment 3, particles from the sample centre starting to be transported).

219

216

220 **3.2. Effect of gravimetric water content**

Figure 3 presents the experimental U^* values as a function of W, as well as RSM regression (contour plot) for

222 moment 3. In order to make Figure 3a clearer, standard deviations were omitted and presented as

supplementary material (Table S2). The RSM regression presented a good fit, with $R^2 = 0.89$ (Table 3). In the

test for the significance of the regression, the rejection of the null hypothesis (p-value < 0.001) implies that at least one of the independent variables contributed significantly to the model.



Figure 3. Threshold friction velocities (U^*) as a function of gravimetric water content (a) and respective response surface (b).

229

226

Table 3. Coefficients of the response surface methodology regression and goodness-of-fit.

| B_0 | B_1 | B_2 | B _{1,1} | B _{1,2} | B _{2,2} | \mathbf{R}^2 | Significance (p-value) |
|--------|--------|--------|-------------------------|-------------------------|-------------------------|----------------|---------------------------|
| 0.4269 | 0.0438 | 0.1156 | -0.019 | 0.022 | -0.009 | 0.89 | < 0.0001 |

231

Figure 3a shows that up to 10 % moisture content the experimental U^* values did not change considerably. 232 233 From 10 % to 20 % moisture content, significant increases were observed for U^* , i.e. 45 % and 25 % for the 234 smallest (d_m of 40 and 125 µm, p-value <0.001) and the largest ($d_m = 6,000 \text{ µm}, \text{ p-value <0.001}$) particles, 235 respectively. This result indicates that erosion of both small and large biochar particles was affected by 236 gravimetric water content, although possibly different mechanisms. The remaining particle size classes tended 237 to maintain similar U^* values, regardless of moisture conditions. However, the response surface applied to the 238 experimental data (Figure 3b), clearly illustrates that the positive effect of moisture content on U^* was more pronounced for small particle size, with an increase of predicted values of U^* from 0.27 to 0.40 m s⁻¹. On the 239 contrary, for coarser particles, the influence of moisture content seemed to be almost negligible, varying 240 between 0.54 and 0.56 m s⁻¹. Indeed, the B_2 coefficient was higher than B_1 (0.12 and 0.04, respectively), which 241 242 reflects a stronger effect of particle size on predicted U^* rather than moisture content. Although the fitting was 243 very accurate, the model tended to overestimate U^* for particles with $d_m = 1,100 \,\mu\text{m}$. In addition it can be argued that the model was heavily weighted towards smaller particle classes ($d_m = 40$ and 125 µm), which are 244 245 relatively close in the experimental range and, therefore, a potential source of bias (Myers et al., 2009).

These results clearly provided important implications of moisture content in biochar applied in the field. Since the smallest particles are those subjected to eventual erosion episodes, it is important to bear in mind the effect of water content to mitigate their detachment. Therefore, U^* of the smallest particles can be increased by keeping *W* at 15 % or higher.

250

3.3. Contributions of adhesive and gravity forces

Table 4 presents, for each particle size, the contributions of adhesive (parameter *b* in Eq. (5) and gravity (parameter *a*) forces, calculated using the experimental data for moment 3. The contribution of gravity (*a*) increased with increasing particle size, showing that the weight gain by water absorption by the particles is more relevant to U^* than adhesive forces. The contribution of adhesive forces (*b*) was more pronounced for small particles (d_m of 40 and 125 µm) rather than larger ones, with *b* values for the former (about 8.5 × 10⁻³) substantially higher than those for the latter (in the range of $2.4 \times 10^{-3} - 4.4 \times 10^{-3}$).

Table 4. Contributions of gravity (*a*) and adhesive forces (*b*) for biochar particle threshold friction velocity (U^*) according to a linear regression model (Eq. (5)).

| d_m (µm) | 40 | 125 | 1,100 | 2,575 | 4,075 | 6,000 |
|------------------------|-----------------------|-----------------------|-----------------------|-----------------------|-----------------------|-----------------------|
| a | 2.46×10 ⁻¹ | 2.32×10 ⁻¹ | 2.86×10^{-1} | 4.05×10^{-1} | 4.55×10^{-1} | 4.94×10 ⁻¹ |
| b | 8.48×10 ⁻³ | 8.53×10 ⁻³ | 2.82×10^{-3} | 2.38×10 ⁻³ | 2.92×10 ⁻³ | 4.45×10 ⁻³ |
| \mathbf{R}^2 | 8.93×10 ⁻¹ | 9.16×10 ⁻¹ | 7.44×10^{-1} | 5.63×10 ⁻¹ | 3.32×10 ⁻¹ | 5.73×10 ⁻¹ |
| Significance (p value) | 4.40×10 ⁻³ | 2.70×10 ⁻³ | 2.70×10 ⁻² | 8.57×10 ⁻² | 2.30×10 ⁻¹ | 8.12×10 ⁻² |

260

The relatively low R^2 values for the larger particle size classes, dm=2,575 µm and above, are primarily related 261 with the non-linear patterns observed in U* for some of the particle size classes, as can be observed in Figure 262 3a. The goodness-of-fit of the three smaller particle size classes, with R^2 values between 0.744 and 0.916, 263 with significant and highly significant regressions, are of the same order as those reported by Zhang et al. 264 (2012), R² values between 0.689 and 0.810. However, the goodness-of-fit for the three larger particle size 265 classes, R² values between 0.573 and 0.332, with non-significant regressions, are lower than those reported by 266 267 Zhang et al. (2012). A subset of data presenting reduced adhesive effect was selected by excluding the 268 smallest particle size classes ($d_m = 40$ and 125 µm). This subset for each of the gravimetric water content series was applied to the modified Bagnold model (Zhang et al., 2012; Eq. (6)), by using air density (ρ_{air}) 269 value of 1.2 kg m⁻³ and an average biochar particle density (ρ_{biochar}) of 450 kg m⁻³ (Shenbagavalli and 270 271 Mahimairaja, 2012). Determined values of A ranged from 0.09 to 0.13 and, under the best fit ($R^2 = 0.87$) it 272 equalled 0.12. As shown in Figure 4, by taking A = 0.12, the modified Bagnold's model accurately predicted 273 the experimental values of U^* for biochar particles that were not affected by adhesive forces. However, this model underestimated U^* for the smaller particle size classes (d_m of 40 and 125 µm), thereby confirming the 274 275 results found by (Zhang et al., 2012) for coal particles with mean diameters smaller than 250 µm.



276

Figure 4. Experimental and calculated U^* of biochar particles using Eq. (6)

278

279 The modified Bagnold's model can be extended to include the contribution of adhesive forces (see section 280 2.3). The resulting Eq. (8) was fitted to the complete experimental data set from this study (including the 281 smaller particle sizes classes) in order to estimate c and d, using the previously determined parameter A =282 0.12. The goodness-of-fit was lower for $W \ge 15$ % (Table 5), most likely caused by strong adhesive forces at those moisture contents, which increased U^* for the smaller particle size classes (as can be seen in Figure 2), 283 thereby decreasing its correlation coefficient. Nevertheless, under the best fit ($R^2 = 0.86$), c and d were 0.87 284 285 and 0.71, respectively. These values were applied to equation 7 to calculate the adhesive effect parameter (β), shown in Figure 5. The calculated β values increased exponentially with decreasing particle size, from 1.7 to 286 12.9 for W = 1 %, and from 61.2 to 452.5 for W = 20 %. 287

- 288
- Table 5. Estimation of c and d coefficients using the modified Bagnold's model (Eq. 8).

| W(%) | 1 | 3 | 6 | 10 | 15 | 20 | |
|----------------|------|------|------|------|------|------|--|
| с | 0.87 | 1.10 | 1.36 | 1.80 | 2.60 | 3.00 | |
| d | 0.71 | 0.70 | 0.68 | 0.74 | 0.87 | 0.87 | |
| \mathbf{R}^2 | 0.87 | 0.79 | 0.79 | 0.80 | 0.45 | 0.39 | |

290





Figure 5. Adhesive effect parameter (β) as a function of gravimetric moisture content and mean particle size (segments between dots provide only a visual connection between calculated β values in the same *W* series).

294

295 4. Discussion

296 In comparison to studies examining the effect of moisture content and particle size on U^* in coal samples, 297 there are some broad similarities, but also important differences. Namely, the studies by Duo-min and Shu-298 tang (1991) and Zhang et al. (2012) showed similar increases in U^* with increase in particle size. However, 299 the particle size range in the present biochar experiment ($<50 \mu m$ to $>6,000 \mu m$) is much greater than in the 300 referred coal experiments ($<45 \mu m$ to $>1,000 \mu m$), reflecting common differences in particle size distributions 301 of woody biochars and coal. Even at large sizes (>6 cm), biochar particles reached U^* values commonly 302 found for sand particles of 0.4 - 0.5 cm (Fécan et al., 1999). This is most likely explained by the much greater 303 density of sand than biochar, both regarding bulk and particle density. Due to lower inter-particle porosity, 304 sandy soils would be subjected to less air entrainment into the soil. In addition, large biochar particles present 305 a platy morphology compared to spheroidal sand particles, possibly resulting in greater shear stress as air 306 passed over and under the large biochar particles, compared to sand particles. This would suggest that even 307 large biochar particles exposed at the soil surface may also become mobile (by creep or saltation mechanisms) 308 under erosive conditions, i.e. dry bare soil on windy days. However, this study was focused exclusively on the 309 response of biochar particles to controlled varying wind conditions, and additional research on the interaction 310 between soil particles and biochar particles, in wind tunnel simulations and in the field, is needed to provide 311 further quantification.

The effect of moisture content also showed a similar pattern for biochar particles as for coal particles (Duomin and Shu-tang, 1991; Zhang et al., 2012). The two smaller particle sizes ($d_m < 125 \mu m$) exhibited increasing

314 adhesive forces with increasing moisture content, with particles $>1,100 \,\mu m$ experiencing 2-3 times lower 315 adhesive effects than particles $\leq 125 \,\mu$ m. The largest particles size class, >4,000 μ m (dm=6,000), appears to 316 show a recurring increase in adhesive effect, although still only approximately half as strong as for the two 317 smaller particle sizes. Possibly, the shape of these largest particles (increased platiness) may have contributed 318 to this effect by increased surface contact. However, the woody feedstock biochar used in this study only had 319 significantly (*p*-value < 0.001) increased U* values for fine particles at moisture contents of 15 % and 20 %, 320 whereas for coal particles used by Zhang et al. (2012) this appeared to occur at 6 - 8 % moisture content. 321 Possibly, this difference can be explained by the larger porosity of biochar particles, i.e. 55 – 70% (Brewer et 322 al., 2014), compared to coal particles, i.e. 6 – 20 % (Zhang et al., 2012). Therefore, water films on biochar 323 particle surfaces would only develop at greater moisture contents than for coal particles. The modified 324 Bagnold model (Zhang et al., 2012) estimated U^* reasonably well for particles $\leq 1,100 \,\mu\text{m}$, but markedly worse estimates were found for particles $\geq 2,575 \,\mu\text{m}$. This implies that, for this type of biochar, the modified 325 326 Bagnold model may be used with confidence to estimate the wind velocity and biochar moisture content at 327 which particles $\leq 1,100 \,\mu\text{m}$ are likely to become airborne. However, for larger particles the model needs to be 328 developed further.

329 These findings suggest to only consider moisture contents of 15 % or greater for recommendations (including 330 certifications) regarding storage, transport, and field application of biochar for the purpose of reducing the 331 propensity of particles becoming air-borne. If further increases in U^* can be reached at moisture contents over 332 20 % requires further research. The smallest particle size considered in this experiment was <50 µm. Many 333 health concerns regarding inhalation of particles focus on respirable suspended particles (<10 µm) or fine 334 particles ($<2.5 \,\mu$ m), or even smaller fractions. Although these particle sizes were part of the smallest particle 335 size class used in this study, it is possible that they were undetected. Follow-up experiments with techniques 336 that can detect particles in this size range are required to provide more insight.

Unlike coal, biochar is made from a wide variety of feedstocks, under a range of pyrolysis conditions,
resulting in a variety of biochar physical properties. A woody feedstock biochar was used in this study as
being representative of many biochars being used in experiments. However, biochar is also produced by
pyrolysis of many other feedstocks, such as poultry litter, wastewater sludge, green waste, and biosolids
(Jeffery et al., 2011), which are known to differ strongly in particle size distributions, particle morphology and

density, (capillary) porosity, etc. (Chia et al., 2015). Therefore, this study's results cannot be directly
extrapolated to other biochars, and further studies are urgently needed to fill this knowledge gap and inform

- 344 policy development. In addition, future studies should also consider effects of saltating particles on threshold
- 345 friction velocities of biochars, and bench and field studies should consider how soil aggregation and
- downward movement of fine particles in soils affect biochar erosion by wind.
- 347

348 5. Conclusions

- In order to identify interactions between wind and biochar, we conducted a wind tunnel study to determine 349 350 threshold friction velocities of a woody feedstock biochar for a range of particle sizes and gravimetric 351 moisture contents. For W < 10 %, fine particles ($d_m < 125 \mu$ m) started to erode at U^* as low as 0.27 m s⁻¹ 352 whereas large particles ($d_m > 6,000 \,\mu\text{m}$) eroded at $0.50 - 0.55 \,\text{m s}^{-1}$, showing that large biochar particles 353 exposed at the soil surface may also become suspended at threshold friction velocities commonly found for 354 sand particles. However, for $W \ge 15$ %, the moisture content presented greater impacts on smaller than larger 355 particles, by promoting their adhesion and increasing the threshold friction velocity. In turn, larger particles 356 exhibited resistance to erodibility due to weight gain by water absorption. These results provide important 357 implications of biochar moisture content to inform its sustainable development and application. Since smaller 358 particles are more susceptible to detachment in field, avoidance of wind erosion can be achieved by keeping 359 W at 15 % or more to increase safe storage, transport and field application of biochar.
- 360

361 5. Acknowledgements

- This work was supported by European Funds through COMPETE and by National Funds through the
 Portuguese Fundação para a Ciência e a Tecnologia (FCT) within projects PEst-C/MAR/LA0017/2013,
 EXPLOCHAR (EXPL/AGR-FOR/0549/2013) and CLICURB (EXCL/AAG-MAA/0383/2012), and the postdoctoral grants of J.H. Amorim (SFRH/BPD/48121/2008) and F.G.A. Verheijen (SFRH/BPD/74108/2010).
- 366

367 **6. References**

- Alfaro, S. C., A. Gaudichet, L. Gomes, M. Maillé, 1997. Modeling the size distribution of a soil aerosol
 produced by sandblasting. J. Geophys. Res, 102, 11239–11249. <u>http://dx.doi.org/10.1029/97JD00403</u>
- Andreae, M.O., Merlet, P., 2001. Emission of trace gases and aerosols from biomass burning. Global
 Biogeochem. Cycles 15, 955-966. <u>http://dx.doi.org/10.1029/2000gb001382</u>
- 372 Bagnold, R.A., 1941. The physics of blown sand and desert dunes. London: Methuen.
- Blackwell, P., Krull, E., Butler, G., Herbert, A., Solaiman, Z., 2010. Effect of banded biochar on dryland
 wheat production and fertiliser use in south-western Australia: an agronomic and economic perspective.
 Soil Research 48, 531-545. <u>http://dx.doi.org/10.1071/SR10014</u>
- Borrego, C., Costa, A.M., Amorim, J.H., Santos, P., Sardo, J., Lopes, M., Miranda, A.I., 2007. Air quality
 impact due to scrap-metal handling on a sea port: A wind tunnel experiment. Atmos. Environ. 41, 63966405. http://dx.doi.org/10.1016/j.atmosenv.2007.01.022
- Boton, M., Azéma, E., Estrada, N., Radjaï, F., Lizcano, A., 2013. Quasistatic rheology and microstructural
 description of sheared granular materials composed of platy particles. Physical Review E 87, 032206.
 <u>http://dx.doi.org/10.1103/PhysRevE.87.032206</u>

- Brewer, C.E., Chuang, V.J., Masiello, C.A., Gonnermann, H., Gao, X., Dugan, B., Driver, L.E., Panzacchi, P.,
 Zygourakis, K., Davies, C.A., 2014. New approaches to measuring biochar density and porosity. Biomass
 Bioenergy 66, 176-185. <u>http://dx.doi.org/10.1016/j.biombioe.2014.03.059</u>
- 385 Buchsbaum, L., 2007. Railroads and shippers clash over coal dust. Coal Age 112, 14-15
- Chen, W., Zhibao, D., Zhenshan, L., Zuotao, Y., 1996. Wind tunnel test of the influence of moisture on the
 erodibility of loessial sandy loam soils by wind. J. Arid Environ. 34, 391-402.
 http://dx.doi.org/10.1006/jare.1996.0119
- Chia, C.H., Downie, A., Munroe, P., 2015. Characteristics of Biochar: Physical and Structural Properties, in:
 Lehmann, J., Joseph, S. (Eds.), Biochar for Environmental Management Science and Technology
 Routledge.
- Choobari, O.A., Zawar-Reza, P., Sturman, A., 2014. The global distribution of mineral dust and its impacts on
 the climate system: A review. Atmospheric Research 138, 152-165.
- 394 <u>http://dx.doi.org/10.1016/j.atmosres.2013.11.007</u>
- Cornelis, W.M., Gabriels, D., 2003. The effect of surface moisture on the entrainment of dune sand by wind:
 an evaluation of selected models. Sedimentology 50, 771-790. <u>http://dx.doi.org/10.1046/j.1365-</u>
 <u>3091.2003.00577.x</u>
- De Capitani, E.M., Algranti, E., Handar, A.M.Z., Altemani, A.M.A., Ferreira, R.G., Barbosa Balthazar, A.,
 Cerqueira, E.M.F.P., Sanae Ota, J., 2007. Wood charcoal and activated carbon dust pneumoconiosis in
 three workers. Am. J. Ind. Med. 50, 191-196. <u>http://dx.doi.org/10.1002/ajim.20418</u>
- 401 Dey, S., 1999. Sediment threshold. Appl. Math. Model. 23, 399-417. <u>http://dx.doi.org/10.1016/S0307-</u>
 402 <u>904X(98)10081-1</u>
- 403 Dong, Z., Liu, X., Wang, H., Wang, X., 2003. Aeolian sand transport: a wind tunnel model. Sediment. Geol.
 404 161, 71-83. http://dx.doi.org/10.1016/S0037-0738(02)00396-2
- 405 Fécan, F., Marticorena, B., Bergametti, G., 1999. Parametrization of the increase of the aeolian erosion
 406 threshold wind friction velocity due to soil moisture for arid and semi-arid areas. Ann. Geophys. 17, 149407 157. <u>http://dx.doi.org/10.1007/s00585-999-0149-7</u>
- Ferreira, A.D., Viegas, D.X., Sousa, A.C.M., 2003. Full-scale measurements for evaluation of coal dust
 release from train wagons with two different shelter covers. Journal of Wind Engineering and Industrial
 Aerodynamics 91, 1271-1283. <u>http://dx.doi.org/10.1016/S0167-6105(03)00077-1</u>
- Genis A, Vulfson L, Ben-Asher J, 2013. Combating wind erosion of sandy soils and crop damage in the
 coastal deserts: Wind tunnel experiments. Aeolian Research 9, 69-73.
- 413 http://dx.doi.org/10.1016/j.aeolia.2012.08.006

- Han, Q., Qu, J., Zhang, K., Zu, R., Niu, Q., Liao, K., 2009. Wind tunnel investigation of the influence of
- surface moisture content on the entrainment and erosion of beach sand by wind using sands from tropical
- 416 humid coastal southern China. Geomorphology 104, 230-237.
- 417 <u>http://dx.doi.org/10.1016/j.geomorph.2008.08.016</u>
- Harrison, R.M., Jones, A.M., Gietl, J., Yin, J., Green, D.C., 2012. Estimation of the contributions of brake
 dust, tire wear, and resuspension to nonexhaust traffic particles derived from atmospheric measurements.
- 420 Environ Sci Technol 46, 6523-6529. <u>http://dx.doi.org/10.1021/es300894r</u>
- Hashizume, M., Ueda, K., Nishiwaki, Y., Michikawa, T., Onozuka, D., 2010. Health effects of Asian dust
 events: a review of the literature. Nihon Eiseigaku Zasshi 65, 413-421.
- He J-J, Cai Q-G, Tang Z-J, 2008. Wind tunnel experimental study on the effect of PAM on soil wind erosion
 control. Environ Monit Assess, 145, 185–193. <u>http://dx.doi.org/10.1007/s10661-007-0028-1</u>
- 425 IPCC, 2013. Climate Change 2013: The Physical Science Basis. Contribution of Working Group I to the Fifth
- Assessment Report of the Intergovernmental Panel on Climate Change, in: Stocker, T.F., D. Qin, D.,
- 427 Plattner, G.-K., Tignor, M., Allen, S.K., Boschung, J., Nauels, A., Xia, Y., Bex, V., Midgley, P.M.
- 428 (Eds.). Cambridge University Press, Cambridge, UK and New York, USA, p. 1535
- Ito, A., Penner, J.E., 2005. Historical emissions of carbonaceous aerosols from biomass and fossil fuel burning
 for the period 1870–2000. Global Biogeochem. Cycles 19, GB2028.
- 431 <u>http://dx.doi.org/10.1029/2004gb002374</u>
- Jacobson, M.Z., 2001. Strong radiative heating due to the mixing state of black carbon in atmosphericaerosols. Nature 409, 695-697.
- 434 <u>http://www.nature.com/nature/journal/v409/n6821/suppinfo/409695a0_S1.html</u>
- Jeffery, S., Verheijen, F.G.A., van der Velde, M., Bastos, A.C., 2011. A quantitative review of the effects of
 biochar application to soils on crop productivity using meta-analysis. Agric., Ecosyst. Environ. 144, 175187. <u>http://dx.doi.org/10.1016/j.agee.2011.08.015</u>
- Karanasiou, A., Moreno, N., Moreno, T., Viana, M., de Leeuw, F., Querol, X., 2012. Health effects from
 Sahara dust episodes in Europe: Literature review and research gaps. Environ. Int. 47, 107-114.
 http://dx.doi.org/10.1016/j.envint.2012.06.012
- 441 Lehmann, J., 2007. Bio-energy in the black. Front. Ecol. Environ. 5, 381-387. <u>http://dx.doi.org/10.1890/1540-</u>
 442 <u>9295(2007)5[381:bitb]2.0.co;2</u>
- Lehmann, J., Gaunt, J., Rondon, M., 2006. Bio-char Sequestration in Terrestrial Ecosystems A Review.
 Mitig Adapt Strat Glob Change 11, 395-419. <u>http://dx.doi.org/10.1007/s11027-005-9006-5</u>
- Lehmann, J., Joseph, S., 2015. Biochar for Environmental Management: An Introduction in: Lehmann, J.,
- Joseph, S. (Eds.), Biochar for Environmental Management Science and Technology Routledge.

- Lin, D.M., Tsai, S.T., 1991. The relation of the threshold velocity of coal dust to its size and humidity. Appl
 Math Mech 12, 513-519. <u>http://dx.doi.org/10.1007/bf02015564</u>
- Myers, R.H., Montgomery, D.C., Anderson-Cook, C., 2009. Response Surface Methodology: Process and
 Product Optimization Using Designed Experiments 3rd ed. Wiley, New York.
- 451 Nguyen, B.T., Lehmann, J., 2009. Black carbon decomposition under varying water regimes. Org. Geochem.
 452 40, 846-853. <u>http://dx.doi.org/10.1016/j.orggeochem.2009.05.004</u>
- 453 Salvi, S.S., Barnes, P.J., 2009. Chronic obstructive pulmonary disease in non-smokers. Lancet 374, 733-743.
 454 <u>http://dx.doi.org/10.1016/s0140-6736(09)61303-9</u>
- 455 Schmidt, H.P., Abiven, S., Kamman, C., Glaser, B., Bucheli, T., Leifield, J., 2012. Guidelines for biochar
 456 production: European biochar certificate version 4.2. Delinat-Institut für Ökologie und Klimafarming,
 457 Arbaz (Switzerland)
- 458 Shenbagavalli, S., Mahimairaja, S., 2012. Production and characterization of biochar from different biological
 459 wastes. International journal of plant, animal and environmental science 1, 197-201
- 460 Terzaghi, K., 1996. Soil mechanics in engineering practice. John Wiley & Sons.
- 461 Toraño, J.A., Rodriguez, R., Diego, I., Rivas, J.M., Pelegry, A., 2007. Influence of the pile shape on wind
 462 erosion CFD emission simulation. Appl. Math. Model. 31, 2487-2502.
 463 http://dx.doi.org/10.1016/j.apm.2006.10.012
- Verheijen, F., Jeffery, S., Bastos, A., Van der Velde, M., Diafas, I., 2010. Biochar application to soils A
 Critical Scientific Review of Effects on Soil Properties, Processes and Functions, Office for the Official
 Publications of the European Communities, p. 149
- Verheijen, F.G.A., Graber, E.R., Ameloot, N., Bastos, A.C., Sohi, S., Knicker, H., 2014. Biochars in soils:
 new insights and emerging research needs Introduction. Eur. J. Soil Sci. 65, 22-27.
 http://dx.doi.org/10.1111/ejss.12127
- Whitman, T., Nicholson, C.F., Torres, D., Lehmann, J., 2011. Climate Change Impact of Biochar Cook Stoves
 in Western Kenyan Farm Households: System Dynamics Model Analysis. Environ. Sci. Technol. 45,
 3687-3694. http://dx.doi.org/10.1021/es103301k
- Zhang, X., Chen, W., Ma, C., Zhan, S., 2012. Modeling the effect of humidity on the threshold friction
 velocity of coal particles. Atmos. Environ. 56, 154-160.
- 475 <u>http://dx.doi.org/10.1016/j.atmosenv.2012.04.015</u>
- Zhang, X., Chen, W., Ma, C., Zhan, S., 2013. Modeling particulate matter emissions during mineral loading
 process under weak wind simulation. Science Total Environment 449, 168-173.
- 478 <u>http://dx.doi.org/10.1016/j.scitotenv.2013.01.050</u>

479

Highlights

- Wind erosion of biochar was assessed through wind tunnel simulations
- Moisture content lower than 10 % does not prevent erosion of fine particles
- Higher moisture increases adhesion of fine particles and weight of large particles
- Minimum 15 % moisture is recommended for reducing wind erosion of biochar

A ALANCE